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Single-particle, particle-pair, and multiparticle dispersion of fluid particles in forced stably stratified turbulence

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The dispersion of fluid particles in statistically stationary stably stratified turbulence is studied by means of direct numerical simulations. Due to anisotropy of the flow, horizontal and vertical dispersion show different behavior. Single-particle dispersion in horizontal direction is similar to that in isotropic turbulence for short times, but shows a long-time growth rate proportional to \( t^{1.2 \pm 0.1} \), larger than the classical linear diffusion limit. In vertical direction, three successive regimes can be identified: a classical \( t^2 \)-regime, a plateau that scales as \( N^{-2} \), and a diffusion limit where dispersion is proportional to \( t \). By forcing the flow and performing long-time simulations, we are able to observe this last regime, which was predicted but not observed before in stratified turbulence. This diffusive regime is caused by molecular diffusion of the active scalar (density). The mean squared separation of particle pairs (relative dispersion) in vertical direction shows two plateaus that are not present in isotropic turbulence. They can be associated with the characteristic layered structure of the flow. In the long-time limit again a linear regime is found as for single-particle dispersion. Pair dispersion in horizontal direction behaves similar to that in isotropic turbulence except for long times. Finally, the study of multiparticle statistics in stably stratified turbulent flows is reported. The evolution of tetrads gives an impression of the shape of particle clouds. It is found that with increasing stratification, the volume of the tetrads decreases, and they become flatter and more elongated. © 2008 American Institute of Physics. [DOI: 10.1063/1.2838593]

I. INTRODUCTION

Dispersion of particles in stratified flows plays an important role in geophysical environments. These particles can be passive or active, like aerosols in the atmospheric boundary layer and micro-organisms in coastal areas and estuaries. Including all biological and physical parameters in modeling particle dispersion is complicated, and as a starting point in this work, the effect of stratification on dispersion of passive fluid particles is considered. Homogeneous stratified turbulent flow is used here, in correspondence with previous fundamental studies. The work that has been carried out on fluid particle dispersion in homogeneous stratified turbulence is to date rather limited. Kimura and Herring1 studied dispersion in decaying stratified turbulence by means of direct numerical simulations (DNS), and Nicolleau and Vassilicos2 used kinematic simulations (KS) to study dispersion in nondecaying stratified turbulence. More recently, Lichtenstein et al.3,4 investigated the influence of nonlinear effects on fluid particle dispersion by comparing the results of DNS simulations of decaying stratified (and rotating) turbulence with those of KS and rapid distortion theory (RDT). Some examples of experimental studies on particle dispersion in stratified turbulence can be found in Pearson et al.5 As opposed to the before-mentioned numerical work, we apply forcing in our DNS of the Boussinesq equations, which makes it possible to study statistically stationary homogeneous stratified turbulence. In this way, we have been able to track particles for sufficiently long times in order to obtain long time series for calculating Lagrangian statistics.

The present study deals with turbulent flows affected by stable stratification, so the average density of the fluid is decreasing with height. Under the influence of strong stable density stratification, two sorts of motion occur simultaneously: propagating internal gravity waves and a nonpropagating nonlinear component connected with quasihorizontal motions.5 The type of flow under consideration is homogeneous and anisotropic. In order to keep the stratified turbulence statistically stationary, energy has to be continuously added in our simulations to account for energy losses due to viscous dissipation. This has the great advantage that the relative importance of stratification effects to turbulence effects is constant in time, as opposed to decaying stratified turbulence, where the influence of stratification increases in time. There is still some discussion how these types of anisotropic flows should be forced, whether forcing should be applied to all three directions or just in the horizontal plane, and at which length scales the turbulence should be forced.7–9 Anisotropy of the flow, and as a result also anisotropy of the dispersion, should evolve naturally from the interaction between the flow and the background density profile and should not be artificially imposed by the forcing method. To this end, different types of forcing have been

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tested. The final choice, which best served our purposes, was forcing only at the largest scales and with equal strength in all three directions.

This study specifically aims to look at fluid particle dispersion. In view of future applications where aggregate formation might play a role, not only the spreading of single particles will be studied but also that of small clusters of particles. For homogeneous isotropic turbulence, a lot of work has been reported on both single-particle and particle-pair dispersion, experimentally but mainly using numerical simulations. \textsuperscript{10–15} Single-particle dispersion is defined here as the mean-squared displacement of a particle from the initial position, and particle-pair dispersion is the mean-squared separation of a pair of particles. Theoretically, the spreading of particles goes like \( t^2 \) for short times (ballistic regime) and is proportional to \( t \) in the long-time diffusion limit. When stable background stratification is present, dispersion in the vertical direction is suppressed. As a result, a plateau is found for single-particle dispersion in the vertical direction, which scales proportional to \( N^{-2} \) with \( N \) the buoyancy frequency. \textsuperscript{1,2} This plateau is reached for intermediate times, around \( t \approx 2 \pi / N \). In the present work, this plateau is found too, but moreover we find a diffusive regime for long times. The effect of stratification on horizontal dispersion is not clear yet, often it is assumed to be similar to dispersion in homogeneous isotropic turbulence. \textsuperscript{1,2,16} Liechtenstein \textit{et al.} \textsuperscript{4} find a long-time behavior that is proportional to \( t^2 \) when using the linear RDT model, whereas their DNS results show a slight tendency toward a diffusive regime. Our work, resulting from long time series of statistically stationary stratified turbulence, shows a clear superdiffusive regime for long-time horizontal dispersion, which is proportional to \( t^{\alpha} \) with \( \alpha = 2.1 \pm 0.1 \).

Attempts to model single-particle dispersion in stratified flows started with the work of Csanady. \textsuperscript{17} More recently, next to the above-mentioned plateau, a theoretical model by Pearson \textit{et al.} \textsuperscript{5} predicts a linear growth of the mean-squared vertical displacement for large times.

For particle-pair dispersion, the theoretical scaling laws for isotropic turbulence depend on the initial separation between the particles. \textsuperscript{18,19} When stratified turbulence is considered, vertical dispersion is suppressed as for single-particle statistics. Nicolleau and Vassilicos \textsuperscript{2} retrieved a plateau again at \( t \approx 2 \pi / N \) and observed the beginning of a second plateau for large times. These two regimes and a final diffusive regime have been retrieved in the present study. Horizontal pair dispersion shows three regimes, as in isotropic turbulence. Initially, the classical ballistic regime is obtained, followed by a larger slope in the intermediate range and again a smaller slope for long times. For long times, the particles are expected to become uncorrelated, resulting in similar long-time dispersion behavior for single particles and particle pairs. Nicolleau and Vassilicos \textsuperscript{2} indeed found this behavior, with a scaling that is proportional to \( t \). Liechtenstein \textit{et al.} \textsuperscript{4} see a much stronger final growth, which according to them is most likely caused by too short integration times.

To study the shape of a cloud of particles, the evolution of a group of particles—four in this work—can be followed. Experimentally, Lagrangian measurements of a cluster of particles can be used to derive the full set of spatial velocity derivatives, useful, for example, to study vorticity dynamics. \textsuperscript{20} Lagrangian multiparticle statistics are described in homogeneous isotropic turbulence by Pumir \textit{et al.} \textsuperscript{21} and Biferale \textit{et al.} \textsuperscript{22} They found that particle clusters with an initial tetrahedral shape were strongly distorted and tended to become elongated, almost coplanar objects. It will be elucidated in this work that the shape of a group of particles is flatter and more elongated in stratified turbulence compared to isotropic turbulence.

The numerical method used in this work is introduced in Sec. II. Next, in Sec. III, the type of flow resulting from numerical simulations of forced stably stratified turbulence will be described and a validation of the forcing method is given. Thereafter, results are discussed in Sec. IV for both horizontal and vertical dispersion of single particles, particle pairs, and tetrads.

II. NUMERICAL METHOD

A. DNS of the Boussinesq equations

The motion of an incompressible fluid in a stably stratified environment is fully described by the Navier–Stokes equations in combination with an equation that imposes the divergence-free constraint and one for the density. After applying the Boussinesq approximation and using the hydrostatic balance \( \partial \rho / \partial z = -(\rho_0 + \tilde{\rho}) g \), they can be written as

\[
\nabla \cdot \mathbf{u} = 0, \tag{1}
\]

\[
\frac{\partial \mathbf{u}}{\partial t} + \mathbf{u} \cdot \nabla \mathbf{u} = -\frac{1}{\rho_0} \nabla \rho' - \frac{\rho'}{\rho_0} g \hat{z} + \mathbf{F}_{\text{u,ext}} + \nu \nabla^2 \mathbf{u}, \tag{2}
\]

\[
\frac{\partial \rho'}{\partial t} + \mathbf{u} \cdot \nabla \rho' = w N^2 \frac{\rho_0}{g} + \kappa \nabla^2 \rho'. \tag{3}
\]

Herein is \( \mathbf{u} = (u, v, w) \) the velocity in the \( x, y, \) and \( z \) direction, respectively, with \( \hat{z} \) pointing upwards. Furthermore, \( \rho \) is the density, \( \rho_0 \) is the pressure, \( g \) is the gravitational acceleration, \( \nu \) is the molecular viscosity, and \( \kappa \) is the scalar diffusivity. \( \mathbf{F}_{\text{u,ext}} \) represents the external forcing. The buoyancy frequency, or Brunt–Väisälä frequency, is defined as \( N^2 = -g/\rho_0 \partial \rho / \partial z \) and the ratio of \( \nu / \kappa = \text{Sc} \) is the Schmidt number. The density \( \rho = \rho_0 + \tilde{\rho} + \tilde{\rho}'(x, y, z, t) \) is split in three components: a typical value \( \rho_0 \) plus a time-independent linear background profile \( \langle \tilde{\rho} \rangle \) plus a fluctuating part \( \langle \rho' \rangle \). The choice of a linear background implies a homogeneous stratification. Fluctuating components are indicated with a prime, and in the following an overbar is used for an averaged quantity. For properties of the flow field, this average is a spatial average over the computational domain, whereas Lagrangian statistics are calculated from ensemble averages over all particles. Note that in taking both averages, we used homogeneity in both horizontal and vertical directions.

The equations of motion are solved using a three-dimensional parallel pseudospectral DNS code (see Ref. \textsuperscript{23} for details). DNS enables us to solve the Navier–Stokes equations exactly at all relevant scales in the flow without making use of any model. The main drawback is that only
relatively low resolutions can be used, thus flows with relatively low Reynolds numbers can be solved due to its high computational costs. The results presented here are derived from simulations with a resolution of $128^3$ to be able to track particles for very long times. However, as a check most cases are studied also at a higher resolution ($256^3$) and they gave similar results for the time range that could be resolved at that resolution. One of these higher-resolution results is included in Fig. 2 as a demonstration. A cubic domain of width $L_0=1$ is used. Periodic boundaries are implemented in all three directions, allowing the use of a Fourier representation of the velocity and scalar field. Time-stepping of the linear viscous term is performed using exact integration, whereas the other terms are treated by a third-order Adams–Bashforth method.

In a precomputation, a divergence-free homogeneous isotropic turbulent velocity field is created using forcing of the flow by injecting energy at the largest scales, equally in all three directions. A general description of the forcing scheme is

$$ F^{n+1}(k) = (1-\alpha)F^n(k) + AR(k)e^{i\phi(k)}, $$

where $(1-\alpha)F^n$ (with $\alpha=[0,1]$) denotes a memory effect of the forcing. $A$ is the forcing amplitude, $R$ is a random value for the forcing amplitude taken from a Gaussian with zero mean and standard deviation 1, and $\phi$ adds a random phase to every forced wavenumber mode. $R$ and $\phi$ have different values for each forced wavenumber and at each forcing time. $F^{n+1}$ and $F^n$ are the forces at forcing times $n+1$ and $n$, respectively. The value of the force is updated every $m$ time step (forcing time $t_f=m\Delta t$) with $m=O(5-10)$. Forcing is only applied to the largest scales of the velocity field with wavenumber modes $0<k\leq 2\pi/k_0$ ($k_0=2\pi/L_0$ the smallest wavenumber), and serves to keep the total kinetic energy statistically stationary. Here $k=||k||$ is the length of the wavenumber vector $\mathbf{k}=(k_x,k_y,k_z)$ in which wavenumber $k_i$, with $i\in\{x,y,z\}$, is a multiple of $k_0$.

The initial conditions of the simulations of forced homogeneously stratified turbulence are a velocity field of homogeneous isotropic turbulence together with a zero-valued scalar fluctuation field. At $t=0$, a linear stable background stratification is switched on by assigning a nonzero value to $N$, which is kept constant throughout the simulation. Four different levels of the background stratification are studied, whereby $N^2$ is varied by a factor of 10 between the different cases. Furthermore, a simulation of homogeneous isotropic turbulence is run in which $N=0$ s$^{-1}$. The forcing method mentioned above is used of which the amplitude $A$ and the forcing time $t_f$ are adapted to keep the kinetic energy statistically stationary. The amplitude $A$ is chosen as high as allowed by the requirement of stationarity, to reach the highest possible Reynolds numbers. Some properties of the different runs are given in Table I.

In order to check whether forcing with equal strength in all three directions is allowed for the resulting anisotropic flow, we also performed a simulation in which only wave-number modes in the range mentioned above and with $k_z=0$ are forced (case horF). In this way, inducing vertical fluid motion by the artificial forcing is avoided; velocity fluctuations in vertical direction are only created via nonlinear interaction with the horizontal velocity components. The relative importance of stratification can be expressed by the vertical Froude number, which is defined here as $Fr_z = u_{rms}/NL_z$, and it gives the ratio of inertial forces to buoyancy forces. The root-mean-squared velocity is given by $u_{rms}^2 = \frac{1}{2}E_{kin}$ and $L_z$ is the vertical integral length scale given by $\int_0^{L_z} R_{wv}(z)\,dz$, with $R_{wv}(z)$ the Eulerian spatial velocity autocorrelation function in vertical direction. The total kinetic energy per unit mass, $E_{kin}$, is calculated as $E_{kin} = \frac{1}{2}(u^2+v^2+w^2) = 2E_h + E_z$. The results in the horizontal directions $x$ and $y$ do not significantly differ (within 3% for the rms velocities and length scales) and therefore in the following only the averaged horizontal values will be used. An impression of the degree of anisotropy of the flow can be derived from the ratios $u_h/\,u_{rms}, u_z/\,u_{rms}$, and $L_z/L_h$. The horizontal velocity $u_h$ is calculated from the total energy $E_h$ in horizontal direction and similarly the vertical velocity $u_z$ stems from the
vertical kinetic energy $E_v$. For the calculation of the horizontal length scale, a similar formula is used as mentioned above for the vertical length scale, thereby taking the averaged value of the components in the $x$ and $y$ directions. The values for $k_{\max}$, with $k_{\max}$ the highest wavenumber resolved by the grid and $\eta$ the Kolmogorov length scale, are included in the table as a measure for the resolution. For a well-resolved simulation, it is required that $k_{\max} \eta > 1$. A measure for the turbulence intensity is $Re_\kappa = 85$ for the initial velocity field. Here the Reynolds number $Re_\kappa = u_{rms}/\nu$ with length scale $\lambda = \sqrt{\left< u''^2 \right> / \left< \partial u'/\partial x \right>^2}$. Most of the results shown in this work are derived from simulations where $Sc=1$. For one value of the stratification (case N100), simulations are run with different Schmidt numbers ($Sc=0.5, 1, 2, 7$) to study the influence of $Sc$ on particle dispersion.

B. Particle tracking

A natural way to describe turbulent dispersion is the Lagrangian frame of reference, in which the observer is moving with the particle. Here we study fluid particles, which are infinitely small fluid elements that exactly follow the flow. Particle trajectories are derived from

$$\frac{dx_p}{dt} = u_p,$$

with $x_p$ the particle position and $u_p = u(x_p)$ its velocity. The velocity at the particle position can be derived from knowledge of the Eulerian velocity field by use of interpolation. Cubic spline interpolation of the velocity field at the particle position is implemented in the code. Next, the particle trajectories are obtained by numerical integration of Eq. (5). Time integration is performed using the same third-order Adams–Bashforth technique as for the Eulerian velocity field.

Velocity and position time series of 16384 particles are collected for about 50 eddy turnover times $T_E=L_x/u_{rms}$. These particles are grouped in triangular pyramid structures, to be able to study both single-particle and particle-pair statistics and the evolution of tetrads. The initial position of one-quarter of the particles is uniformly spread over the computational domain. The other particles are initially located at a fixed separation (about $\frac{5}{2} \eta$ for cases N1, N10, and N100) in the $x$, $y$, and $z$ directions from the reference particles. For two values of the background stratification (cases N0 and N100), the influence of the initial particle separation on particle-pair and multiparticle statistics is also studied. In these cases, 81920 particles are tracked with initial separations of about $\frac{5}{2}, \frac{5}{2}, \frac{5}{2}, 6,$ and $15$ (in units of $\eta$) in all three directions. Particles are released when the flow has reached a quasi-stationary state.

III. FLOW STRUCTURE IN FORCED STRATIFIED TURBULENCE

A. Validation of the forcing method

An important question in this work is whether it is possible to maintain a state of statistically stationary stably stratified turbulence by applying artificial forcing. If energy transfer to smaller scales is modified or an inverse cascade is present in the flow, forcing of the large scales could lead to accumulation of energy at these scales and eventually to a collapse of the simulation. We do find that the type of forcing described in Sec. II A with equal strength in all three directions results in a quasi-stationary state. Checking stationarity is done by looking both at kinetic energy in all three directions and at the velocity derivative skewness $S_3$. The velocity derivative skewness is defined as

$$S_{3,z} = \frac{\left< (\partial u'/\partial z)^3 \right>}{\left< (\partial u'/\partial z)^2 \right>},$$

for the vertical component, and in an analogous way for the horizontal component, $S_{3,h}$. Time series of $S_{3,h}$ and $S_{3,z}$ are shown in Fig. 1 and their time-averaged values are given in Table II. The value of these quantities for case N0 ($-0.49$) is

![FIG. 1. Horizontal (a) and vertical (b) velocity derivative skewness as function of time for three values of the stratification: N0 ($\cdots$), N10 ($\cdots$), N1000 ($\cdots$). For clarity, cases N1 and N10 are left out. They show similar quasi-steady behavior around values of $-0.52$ and $-0.37$ (horizontal) and $-0.43$ and 0.02 (vertical), respectively.](image-url)
consistent with values found in studies of homogeneous isotropic turbulence. In isotropic turbulence, a negative value of $S_3$ points to a forward energy cascade. Whether this relation holds for stratified turbulence is unknown. For all cases studied in this work, the skewness $S_3$ fluctuates around constant values, demonstrating the statistically stationary state reached in the simulations. With increasing stratification, the value of $S_{3,h}$ first becomes more negative and then goes toward zero, whereas $S_{3,z}$ first increases—even becomes positive—and then decreases also toward zero. This trend is also visible in Fig. 1 of the work by Kimura and Herring. The behavior of the skewness for moderate stratification [case N1 (not shown) and mainly case N10] might be explained by a dominance of the effect of anisotropy of the flow on nonlinear energy transfer. For strong stratification [cases N100 (not shown) and N1000], the velocity derivative skewness goes toward zero, indicating suppressed nonlinear energy transfer from small to large wavenumbers.

Two convincing indications for the existence of a forward energy cascade in forced stably stratified turbulence are found. No accumulation of energy takes place at the largest scales in the flow, which would give a blowup of the total kinetic energy. Moreover, a test in which forcing is applied around $k=8k_0$ resulted in a flow with only a little amount of energy at the largest scales (smallest wavenumbers $k$), displaying only small-scale flow structures. Because of the presence of this clear forward energy cascade, large-scale forcing as applied in this study is appropriate.

A big advantage of performing forced stratified turbulence simulations for dispersion studies is the following. By keeping the energy distribution over the length scales in the flow statistically stationary, the turbulence level is kept constant and hence the relative importance of the background stratification with respect to turbulence (quantified by the Richardson number, see Sec. III B) remains constant. This is not the case for decaying stratified turbulence. The energy spectra of decaying stratified turbulence simulations progressively steepen in the course of time due to dissipation at the smallest scales, thereby increasing the relative importance of the background stratification.

In Fig. 2, the horizontal and vertical kinetic energy are shown as a function of time for a stratification with $N=0.98$ $s^{-1}$. Three runs are shown, one with purely horizontal forcing (case horF) and two with 3D forcing (cases N100 and lowF). The difference between these last two runs is the amount of energy that is added to the flow by the forcing (total kinetic energy differs by a factor of about 2). Cases N10 and N100, where the strongest competition is to be expected between turbulence generation by forcing on the one hand and turbulence suppression due to stratification on the other hand, were most difficult to get completely stationary. A slight increase in time of the averaged horizontal kinetic energy is found (for example, solid line in Fig. 2). A completely flat horizontal kinetic energy level is only achieved when the applied forcing amplitude was small, and the flow could hardly be called turbulent. In terms of kinetic energy, for case lowF it follows from the energy spectra that only approximately 1% (compared to 10% for isotropic turbulence and 3–4% for case N100) of the energy is found at $k/k_0>10$, so almost all small-scale motion has then vanished. The slight increase in the kinetic energy is probably caused by the excitation of internal gravity waves, as discussed by Lindborg and Brethouwer. However, the particle dispersion (mean-squared displacement) in both high (case N100) and low (case lowF) energetic stratified flows is very similar. In particular, the scaling behavior in the different time regimes shows few differences with energy level. For vertical dispersion, this is shown in Fig. 2(b), and also for horizontal dispersion the slopes for both short times and long times are approximately the same for both runs. The results shown in this work are derived from the high energetic case in which the horizontal kinetic energy increases slightly. This choice is based on the following argument. For comparison of the cases N10 and N100 with cases N1 and N1000, where
it was easy to obtain quasistationarity, a comparable amount of small-scale motion is preferred, thereby avoiding a too short range of length scales in the flow.

Furthermore, (an)isotropy of the forcing is tested, because the resulting flow is anisotropic for cases with \( N \neq 0 \). Applying equal forcing in all three directions or purely forcing the horizontal wavenumber modes leads to different flow configurations, with the main differences being the ratio between the vertical and horizontal kinetic energy (see Fig. 2(a)) and the ratio \( L_z/L_h \) (0.32 for 3D forcing, 0.09 for horizontal forcing). Dispersion statistics, however, are qualitatively very similar despite differences in the large-scale flow structures. For the vertical direction, dispersion statistics are shown in Fig. 2(b) for both purely horizontal and three-dimensional forcing (both high and low energetic), as well as for a simulation performed with 3D forcing at higher resolution (256\(^3\)). For horizontal dispersion (not shown), the differences between the two forcing methods (2D and 3D) are hardly visible and the slopes are exactly the same. A detailed description of the results follows in Sec. IV A; for now only the strong resemblance between the two forcing methods is of importance. Horizontal forcing cannot be applied to the isotropic case N0 and the weakly stratified case N1. In order to make all simulations comparable, the same forcing type, and thus 3D, is chosen.

**B. Flow structure and density profile**

To get an idea of the structures in the flow, in Fig. 3 plots of isovorticity surfaces for different stratification levels are shown. The isovorticity surface of case N1 resembles those found for 3D isotropic turbulence; the structures are small and directed in all three directions. The strong stratification (case N100) shows the “well-known” pancake-like structures, whereas for the weaker stratification (case N10) traces of small-scale structures can be seen in addition to the pancakes. The pancake-like structures are related to regions of strong vertical shear.\(^1\) In Fig. 4, the horizontally averaged square of the vertical shear, defined as

\[
\Sigma^2 = \left( \left( \frac{\partial u}{\partial z} \right)^2 + \left( \frac{\partial v}{\partial z} \right)^2 \right)_h,
\]

is plotted as a function of \( z \) for case N100. Alternating regions of high and low vertical shear demonstrate the layered structure. According to Brethouwer et al.,\(^26\) the dynamics of turbulence affected by stable stratification can be classified into two distinct regimes. As in the work by Billant and Chomaz,\(^27\) they introduce a parameter \( R = \text{Re}_h \text{Fr}_h^2 \) with \( \text{Re}_h = u_{\text{rms}} L_h / \nu \) the horizontal Reynolds number and \( \text{Fr}_h = u_{\text{rms}} / L_h N \) the horizontal Froude number. For \( R < 1 \), large smooth horizontal layers can be observed and energy is dissipated mainly by vertical shearing of the large-scale pancake-like structures. When \( R > 1 \), large quasi-horizontal layers are still noticeable but at the same time small-scale turbulent-like motions are present and dissipation takes place predominantly at these smallest scales. For the cases studied in the present work, this classification would separate cases N1 and N10 (\( R > 1 \)) from cases N100 and N1000 (\( R < 1 \)).

For case N1, the effect of stratification is only weak, but for case N10 in Fig. 3 indeed large-scale horizontal structures can be already identified in combination with small-scale fluctuations. According to Brethouwer et al.,\(^26\) these small-scale motions are most likely due to Kelvin–Helmholtz-type instabilities generated by shear between different layers. The fact that case N10 reveals both large-scale horizontal structures and appreciable small-scale fluctuations, thus representing highly anisotropic turbulent flow, explains the abnormal values of the skewness (\( S_{\Sigma^2} \approx -0.72 \) and \( S_{\Sigma^2} \approx +0.13 \); see Sec. III A) and the flatness (see below). The influence of large-scale dissipation for cases N100 and N1000, resulting in reduced transport of energy through a cascading process, becomes clear also from the values of the velocity derivative skewness (see Fig. 1).

In a similar fashion as for the skewness, which gives a measure of the asymmetry of a probability density function...
file. There the stable stratification is so strong that turbulent fluctuations are hardly visible. For moderate stratification [Fig. 5(a)], much more overturning and mixing between different layers can be seen. To quantify the amount of mixing, a local gradient Richardson number can be defined as

\[ R_{ig} = \frac{N^2}{\left(\frac{\partial u_h'}{\partial z}\right)^2} \]

which is a ratio of restoring buoyancy forces and turbulence producing shear forces. Here \( u_h' \) is the horizontal component of the fluctuating velocity field at a single moment in time. For steady mean shear flows of an inviscid stably stratified fluid, the Miles–Howard criterion states that the flow can only be unstable and turbulent when \( R_{ig} < R_{ig,c} = 0.25 \). For practical flow situations different values are found, all of order \( O(1) \). Probability density functions of the values of \( R_{ig} \) in the whole computational domain at a single time step are shown for different values of the stratification in Fig. 6, including the reference value 0.25. For case N1, it can be seen that in large parts of the domain, shear effects are dominant \((R_{ig} < 1\) at about 60% of the gridpoints), so a lot of overturning is possible. This picture changes considerably with increasing stratification. For case N10, about 25% of the gridpoints has \( R_{ig} < 1 \), and for case N100 this is \( =0\%\); almost all mixing has vanished there.

Wavelike motion is present in both the velocity and the scalar field. It can be seen in most quantities (Eulerian and}

![Image](https://via.placeholder.com/150)

*(a) (b) (c)*

Fig. 5. Vertical cross sections of the density profile. For case N1 (a) mixing is clearly visible, whereas case N100 (b) only shows slight fluctuations upon the linear background profile.

![Image](https://via.placeholder.com/150)

Fig. 6. Probability density functions of the gradient Richardson number for cases N1, N10, and N100, respectively. The theoretical critical value of \( R_{ig,c} = 0.25 \) is also shown in the graphs (dashed line). With increasing stratification, less turbulent overturning can take place.
Lagrangian, mainly in the vertical) for which the evolution in time is studied, such as kinetic and potential energy, length scales, autocorrelation functions, and particle dispersion; see, for example, the oscillations around the plateau in Fig. 8(b). The frequency of these waves corresponds to the buoyancy frequency, their fundamental frequency $f_0 \approx N$.

The energy distribution between different length scales in the flow is different in stratified turbulence compared to isotropic turbulence. This becomes clear from the energy spectra plotted in Fig. 7. They are calculated as

$$E_h(k) = \sum_{k/k_0 \leq k < k + 1/k_0} \frac{1}{2} \left( |\tilde{u}(k)|^2 + |\tilde{v}(k)|^2 \right),$$  \(10\)

$$E_v(k) = \sum_{k/k_0 \leq k < k + 1/k_0} \frac{1}{2} |\tilde{w}(k)|^2$$  \(11\)

for horizontal and vertical directions, respectively. For the isotropic case, a small region with inertial range scaling ($k^{-5/3}$) can be seen. The influence of stratification on the flow for the weakest stratified case N1 is rather small, as its spectra have almost the same shape as in isotropic turbulence (case N0). All three stronger stratified cases show qualitatively the same behavior. Both in horizontal and vertical direction, the spectrum becomes steeper for stratified turbulence, so less energy is present at the smallest scales. Comparing the horizontal and vertical directions [see Figs. 7(a) and 7(b), respectively], at the largest scales much more energy can be found in the horizontal velocity component than in the vertical velocity component. The horizontal spectrum, however, falls off faster at higher wavenumbers.

The flow is highly anisotropic for the strongest stratified cases, as can be deduced, for example, from the values of the ratios $u_h/u_{rms}$, $u_v/u_{rms}$, and $L_h/L_0$ in Table I. The vertical length scale $L_v$ gives a measure of the layer thickness and the horizontal length scale $L_h$ can be interpreted as the width of the energy-containing eddies. The ratio $L_v/L_h$ decreases with increasing stratification because both $L_v$ decreases and $L_h$ increases. To check whether the growth of the horizontal length scales is not limited by the size of the domain or becomes too large to be able to use periodic boundaries ($L_h$ should be smaller than $0.5L_0$), a simulation is performed with horizontal domain size $2L_0$. The results of this simulation do not show marked differences in both $L_h$ and the dispersion behavior of fluid particles.

IV. PARTICLE DISPERSION IN FORCED STRATIFIED TURBULENCE

A. Single-particle dispersion

The dispersion—or mean-squared displacement—of particles is given by Taylor’s equation,

$$\langle (X(t) - X(0))^2 \rangle = 2u_{rms}^2 \int_0^t (t - \tau) R_L(\tau) d\tau$$  \(12\)

under the assumption that the Lagrangian and Eulerian rms velocities are the same ($u_{rms}^2 = \frac{1}{2} E_{kin}$) and that the flow is homogeneous and stationary. $R_L(\tau)$ is the Lagrangian velocity autocorrelation function, which is only a function of the time separation $\tau = t - t'$,

$$R_L(\tau) = \frac{u_h(t')u_h(t)}{u_{rms}^2}.$$  \(13\)

Using the known properties of the autocorrelation function, $R_L(0) = 1$ and $T_L = \int_0^\infty R_L(\tau) d\tau$ with $T_L$ the Lagrangian time scale, the following relations can be derived for dispersion in homogeneous isotropic turbulence:

$$\langle (X(t) - X(0))^2 \rangle = u_{rms}^2 t^2, \quad t \to 0,$$  \(14\)

$$\langle (X(t) - X(0))^2 \rangle = 2u_{rms}^2 T_L t, \quad t \to \infty$$  \(15\)

(see standard textbooks on turbulence for a derivation of the above equations, for example, Refs. 24 and 29). In the following, $T_L$ for case N0 will be used as a typical Lagrangian time scale. For case N0, the ratio of the Lagrangian time scale and the Eulerian time scale $T_E = T_L/u_{rms}$ is $T_L/T_E = 0.7$, consistent with values found in literature. This ratio decreases with increasing stratification because the typical time scale of the flow $T_E$ increases. For stratified turbulence, several authors (see, for example, Refs. 2, 5, and 30) provided evidence of a plateau for vertical dispersion for $t \geq \frac{3\pi}{2}$. This plateau scales as
Furthermore, the choice of the rms velocity in Eq. (16) differs; some take the overall value as defined in Sec. II A, while others use only the component in vertical direction. We use the vertical component given by \( \bar{w}^2 = 2E_z \) with \( E_z \) the kinetic energy in vertical direction. In the following, horizontal dispersion \( \frac{1}{2} \left( (X(t) - X(0))^2 + (Y(t) - Y(0))^2 \right) \) will be denoted by \( \delta h^2 \) and for the vertical direction \( \delta z^2 = \bar{Z}(t) - Z(0))^2 \) will be used.

As described in the Introduction, horizontal dispersion in stratified turbulence is often assumed to be similar to that in homogeneous isotropic turbulence. Only recently, some first doubts about this assumption have been put forward by Liechtenstein et al. From their work, however, it cannot be deduced whether the superdiffusive long-time regime is inherent to the type of flow under consideration or that it is a result of too short integration times.

Single-particle dispersion in horizontal and vertical directions in stationary stratified turbulence is shown in Fig. 8 for five different values of the density stratification. For isotropic turbulence (case N0), we do retrieve the classical regimes of Eqs. (14) and (15), as well as for horizontal dispersion in relatively weak stratification (case N1). For strong stratification, the plateau with its accompanying scaling is found for vertical dispersion. When the horizontal and vertical axes of Fig. 8(b) are rescaled to \( t/\pi \) and \( \delta z^2 N^2/\bar{w}^2 \), respectively, the onset of the plateaus collapse, as can be seen in Fig. 9(a). For times up to about \( t/T_L = O(10) \), these plots resemble the results of Refs. 1 and 2.

For longer times a new regime can be identified, which becomes available by tracking the particles for sufficiently long times in a quasi-stationary stably stratified flow. Dispersion in vertical direction starts to increase again and is proportional to \( t \), which is a clear indication of a diffusion process. This diffusion of fluid particles away from the original equilibrium position is caused by molecular diffusion of the active scalar (density), which we checked by changing the Schmidt number. The effect of the Schmidt number on dispersion in vertical direction is shown in Fig. 9(b). With increasing \( Sc \), or similarly with decreasing molecular diffusion of the density, vertical dispersion indeed reduces; the diffusive regime sets in at later times. In the limit of \( Sc \to \infty \), only the plateau would be found. The diffusive regime was already predicted by the model of Pearson et al. From their work, it can be deduced that the time scale for the fluid elements to change their density (\( \tau_{trans} \)) is proportional to the Schmidt number, \( \tau_{trans} \propto Sc \). Here we choose this starting time of the diffusive regime as the point at which vertical single-particle dispersion starts to deviate from the results for case \( Sc = 7 \). It is found that these starting times \( \tau_{trans} \) double when \( Sc \) doubles from 0.5 to 1 and from 1 to 2, thereby confirming in this range of Schmidt numbers the predictions made by Pearson et al. This scaling regime has not been observed in decaying stratified turbulence.
The following observations are made for vertical dispersion. In the early stages, particles move away from their initial position with their local velocity. This results in the ballistic $t^2$ regime. Next their vertical displacement is restricted to a thin layer. Particles perform wavelike motion that is also indirectly visible from the continuous increase and decrease of the vertical mean-squared displacement around an averaged plateau level [see Fig. 8(b)]. During this stage, particles continuously convert vertical kinetic energy into potential energy and vice versa.\(^\text{16}\) In case the particles are released in a thin horizontal plane instead of as a homogeneous random distribution over the total domain, particle dispersion denotes the growth of the layer thickness. It is found in a separate simulation that particles released in this way will remain in a thin horizontal layer. For long times finally, due to the exchange of density between elements in the fluid by molecular processes, the actual equilibrium height of each individual particle slowly starts to deviate from its initial equilibrium height. Particles thus forget their initial positions, they obtain new equilibrium levels, and the resulting behavior is diffusive in the long-time limit.

The origin of the diffusive regimes found for dispersion in isotropic turbulence and in stratified turbulence is fundamentally different. In isotropic turbulence, the motion of a particle is purely advective. The linear dispersion behavior for long times is therefore a random-walk-like statistical property.\(^\text{24}\) On the other hand, the diffusive regime for vertical single-particle dispersion in the case of strong stratification is caused by molecular diffusion of the density of the individual fluid particles. Without molecular diffusion, no long-time diffusion limit would be present.

The tool of tracking fluid elements can be used to study flow properties. As shown by Ref. 31, both particle trajectories (Lagrangian) and the diapycnal flux (Eulerian) can be used to derive the eddy diffusivity coefficient when studying mixing of a stably stratified flow. The diffusive regime for vertical dispersion in strongly stratified turbulence found in the present work gives a signature of irreversible mixing due to the diapycnal flux.

The quantitative differences in the long-time behavior of vertical dispersion between 2D and 3D forcing (Fig. 2) can be explained by differences in the flow dynamics. When forcing purely horizontally, mainly the vertical length scale is smaller than in the 3D-forced case. By rescaling the time axis with a large-scale advective time scale based on $L_v$ and $w'$ instead of scaling with $2\pi/N$, the long-time behavior becomes quite similar for both cases.

Horizontal dispersion is enhanced by stratification, as can be seen in Fig. 8(a). For long times, a clear superdiffusive regime is found that is proportional to $t^\alpha$, with $\alpha=2.1 \pm 0.1$. The main cause of the presence of a superdiffusive regime is the effect of vertical shear. Stably stratified turbulence shows quasi-two-dimensional behavior. In the horizontal plane, large-scale vortical structures exist with large eddy turnover time scales. These horizontal structures are not fully decoupled in the vertical direction and strong shearing occurs between vortices in different layers. The vertical range of displacement of fluid particles is limited; in case N100, for example, a fluid particle traverses about 1/20 of the domain height. Locally the flow resembles a shear flow (see Fig. 4), and this shear causes superdiffusive dispersion behavior. Horizontal dispersion in flows with a constant mean vertical velocity gradient (linear shear), for example, is proportional $t^3$.\(^\text{32}\) In the present study, we do not have pure shear flow only. Tests in which random-walk motion is carried out in profiles that resemble parts of the real vertical mean shear profile found in stratified turbulence (see Fig. 4) show the following. In a region of uniform flow [classical picture of a random walk, Fig. 10(a)], the long-time dispersion behaves like $t$; in linear shear flow [Fig. 10(b)], this gives $t^3$; and in the profile depicted in Fig. 10(c), the long-time dispersion limit scales proportional to $t^2$. It is therefore anticipated that the horizontal mean-square displacement in homogeneous stratified turbulence should scale like $t^\alpha$, with $1<\alpha<3$.

The influence of the Schmidt number on horizontal dispersion is negligible, which was expected because of the relative unimportance of density diffusion compared to advection.

**B. Particle-pair dispersion**

In view of future applications, in which aggregate formation (relevant for dispersion of micro-organisms) plays a role, not only single-particle statistics but also particle-pair statistics are of importance. For homogeneous isotropic turbulence, a couple of regimes can be identified for relative dispersion.\(^\text{18,19}\) For short times, the mean-squared separation between two fluid elements grows either exponentially or like $t^2$. When the interparticle separation distance falls in the inertial range, this growth behaves as $t^3$, according to Richardson.\(^\text{33}\) For long times, a diffusion limit exists similar to single-particle dispersion. Which regimes are passed through depends on the initial separation of the particles, only the final $t$ regime is universal since at long times particles become uncorrelated, independent of their initial separation $\Delta_0$. The $t^3$ regime is derived theoretically for high Reynolds number flows with a clear inertial range.\(^\text{34}\)

The separation behavior for homogeneous isotropic turbulence (case N0) can be seen in Fig. 11 for an initial separation of about $\eta$. A $t^2$ regime is expected initially and is indeed visible. Next follows a region that is close to $t^3$, though because of our relatively low Reynolds number this is a necessary transition from the $t^2$ regime to the $t$ regime of
which the levels are determined by the flow instead of real Richardson dispersion. Moreover, this intermediate range with scaling close to $t^3$ is contaminated by the dissipative range because not all particles start separating at the same time. The pair dispersion is therefore a mixture of ballistic $t^2$ and $t^3$ behavior. This effect is demonstrated by Boffetta and Sokolov and Biferale et al. who made use of the method of doubling times as introduced by Artale et al. Finally, the long-term diffusion limit is clearly visible.

Figures 11 and 12 show pair dispersion in vertical ($\xi_z^2(x, t, \mathcal{L})$) and horizontal ($\xi_x^2(x, t, \mathcal{L})$) directions as a function of time, where superscripts 1 and 2 denote the two particles of a pair. These results are derived from pairs with an initial separation in directions perpendicular to the plotted dispersion direction, for example an initial separation in the horizontal plane for $\xi_x^2$. Results of pairs initially separated along the plotted direction are very similar, except for vertical cases in which $\Delta_0$ in the $z$ direction is larger than the levels of the first plateaus. Figure 11(a) shows that also in a stratified environment vertical separation starts with the classical $t^2$ regime. Next a plateau is found, starting around $t = 2\pi/N$. At later times, an increase to a second plateau (indicated with an arrow) can be seen, at least for the strongest stratifications. The time at which the increase to this second plateau starts is of the order of the largest time scales in the flow ($T_2$). The final regime is again the linear diffusion limit. Nicolleau and Vassilicos already predicted the existence of the second plateau, although their kinematic simulation was not carried out long enough to see more than just the beginning of this plateau. The transition from the first to the second plateau coincides with the intermediate $t^3$ regime for horizontal pair dispersion, when particles start moving apart in the horizontal. Although separation in vertical direction is limited, particles can freely disperse in horizontal direction. After some time, the horizontal separation is large enough for the two particles to become uncorrelated. The second plateau can therefore be associated with the plateau found for single-particle dispersion; the level of the second plateau is twice the level of the plateau for the dispersion of a single particle. The influence of initial separation is shown in Fig. 11(b). In general, the behavior is the same for different $\Delta_0$. Especially for long times, its influence is rather small; a linear, uncorrelated behavior is seen at almost the same level. For very large initial separations in the horizontal plane, it is expected to see only one plateau, namely the second one. Already initially particles would reside in different structures within the flow, so there is no argument left for the existence of the first plateau. Indeed, our largest initial separation ($15\eta$) is much smaller than the horizontal extent of the dominant flow structures ($L_h$), and the first plateau is still present.

For horizontal pair dispersion, the effect of stratification...
is less pronounced than for vertical pair dispersion, although the behavior is clearly altered compared to that in isotropic turbulence. As can be seen in Fig. 12(a), the start of the intermediate range is shifted toward larger times and the long-term behavior of case N10 in particular differs from the isotropic case. For long times, the particles forming a pair become uncorrelated. This would lead to the same dispersion behavior as for single particles, which is superdiffusive. Contrary to single-particle dispersion, here the graphs tend toward a linear slope or slightly steeper for long times, except for case N10, where the slope remains $O(t^{2.5})$. Longer integration times are needed to draw a final conclusion about the long-term behavior of horizontal particle-pair dispersion. The onset of the long-time dispersion behavior shifts toward larger times for pair dispersion compared to single-particle dispersion, as can be illustrated by the shift in time between the related plateaus for vertical single-particle dispersion and vertical particle-pair dispersion. The long-time horizontal particle-pair dispersion with possible superdiffusive behavior is not captured in the present simulations.

When looking at the influence of initial separation (Fig. 12(b)), it can be seen that also in the intermediate range stratification influences horizontal pair dispersion, at least for case N1000. For the largest initial separations, a small bump is found in the intermediate range, around $t/T_L$ = 5. This time coincides with the time at which the strongly stratified cases start to deviate from the isotropic case for single-particle dispersion in horizontal direction.

C. Multiparticle statistics

The evolution of a cluster of four particles is studied to get an idea of the shape dynamics of clouds of particles in stably stratified turbulence. Under the action of the flow field, the cloud deforms, and in this way the shape of the cloud can also be used to characterize the flow field. The analysis of multiparticle statistics in this work is conducted along the same lines as in previous work on homogeneous isotropic turbulence.21,22 In order to characterize the shape dynamics for a set of $M=4$ particles, the following set of vectors is introduced:38

$$r_0 = \frac{(x_{p,1} + x_{p,2} + x_{p,3} + x_{p,4})}{2},$$

$$r_1 = \frac{(x_{p,2} - x_{p,1})}{\sqrt{2}},$$

$$r_2 = \frac{(2x_{p,3} - x_{p,2} - x_{p,1})}{\sqrt{6}},$$

$$r_3 = \frac{(3x_{p,4} - x_{p,3} - x_{p,2} - x_{p,1})}{\sqrt{12}},$$

with $x_{p,i}$ ($i=1, M$) the particle positions at the vertices of initially regular (but not isotropic) tetrahedra. $r_0$ defines the center of mass of the cluster, which has no influence on the cluster shape because of the homogeneity of the flow and the uniform initial particle distribution. Vectors $r_1$ to $r_3$ form the columns of a square matrix $r$ from which a moment of inertia tensor $I = \mathbf{rr}^T$ can be defined. The radius of gyration $R^2 = \sum_{i=1}^{M} r_i^T = g_1 + g_2 + g_3$ measures the spatial extent of the swarm of particles. The shape of the cluster of particles can be derived from the eigenvalues $g_i$ ($g_1 \gg g_2 \gg g_3$) of matrix $I$. Often these eigenvalues are given as the ratio $I_i = g_i/R^2$ (obviously $I_1 + I_2 + I_3 = 1$). For the case $M=4$, $g_1 = g_2 = g_3$ corresponds to an isotropic object. When $g_1 = g_2 \gg g_3$, the object is flattened in one direction and has a pancake-like shape and $g_1 \gg g_2, g_3$ represents a needlelike object. Convenient ways to describe the overall shape of an object consist of monitoring $I_2$ or the volume $V = |\det(r)|$ of the object.21 Using Monte Carlo simulations, Pumir et al.21 and Biferale et al.22 derived the values $(I_2)_{GA} = 0.222$ and $(I_2)_{GA} = 0.03$ for an isotropic Gaussian particle distribution, from which it follows that $(I_2)_{GA} \approx 0.748$. Instead of looking only at mean values, a good description of the type of structures occurring in a certain flow is also given by the probability density function (pdf) of the eigenvalues $I_i$.

The procedure described above is first applied to case N0, homogeneous isotropic turbulence. The results from that case can be compared to previous work by Pumir et al.21 and Biferale et al.22 and they serve as a reference when studying the shape evolution of particle clouds in stratified turbulence. The results correspond nicely with Fig. 1 in Pumir et al.21 and Fig. 2 in Biferale et al.22. In Fig. 13, the volume $V$ and scaled eigenvalues $I_i$ are shown as a function of time for different initial cloud sizes. Time is scaled by the Kolmogorov time $\tau_\eta$ for comparison with literature.21,22 The ratio of $\tau_\eta/T_L$ is $O(10^{-1})$. For each different case shown in the figure, the results are ensemble averages over 4096 tetrads. The volume of the tetrahedra grows in time, where the start of the growth occurs slightly earlier for larger $\Delta_0$. In the intermediate range, no convincing $t^{3/2}$ regime (Richardson) is found. This is due to the same reasons as mentioned for particle-pair dispersion in horizontal direction. The long-time dispersion behavior shifts toward larger times for pair dispersion compared to single-particle dispersion, here the graphs tend to a linear slope or slightly steeper for long times, except for case N10, where the slope remains $O(t^{2.5})$. Longer integration times are needed to draw a final conclusion about the long-term behavior of horizontal particle-pair dispersion. The onset of the long-time dispersion behavior shifts toward larger times for pair dispersion compared to single-particle dispersion, as can be illustrated by the shift in time between the related plateaus for vertical single-particle dispersion and vertical particle-pair dispersion. The long-time horizontal particle-pair dispersion with possible superdiffusive behavior is not captured in the present simulations.

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$$r_3 = \frac{(3x_{p,4} - x_{p,3} - x_{p,2} - x_{p,1})}{\sqrt{12}}.$$
separation statistics (see Sec. IV B). For long times, diffusive growth can be seen, related with the uncorrelated motion of the particles that form a tetrahedron. Our results cover a much larger time range than those of Pumir et al. and Biferale et al., making it possible to find this long-time diffusive behavior, not only for the volume but also for the unscaled eigenvalues $g$, (not shown). Looking at the eigenvalues, it can be seen that initially $I_2$ and $I_3$ decrease and $I_1$ increases. The maximum distortion is reached around $t/\tau=O(10)$; the ratio of the eigenvalues $I_1:I_2:I_3$ is then approximately $10^0:10^{-1}:10^{-2}$ (exact values depend on initial separation). This moment falls in the intermediate range, when the growth of the volume is strongest. Geometrically it means that the tetrahedra change from their initial regular shape to more pancake-like and even needlelike structures. They are strongly elongated. For long times, it is expected that smaller tetrahedra reach the peak eigenvalues at later times than larger ones. For the smallest three $\Delta_0$ this is indeed found, however for larger $\Delta_0$ ($6.7, 15.\eta$) the instants of time of maximum distortion increase again. Of course, at a single moment in time all kinds of shapes are possible. The graph shown so far only gives an idea of the mean shape of the objects in the flow. Plots of the probability density function give a good picture of the types of structures occurring at different times during the evolution. For an initial cloud size of the order of the Kolmogorov size, the pdfs of $I_1$ to $I_3$ are plotted on the left-hand side of Fig. 14. This is done for three different times: one shortly after the release of the particles in the flow, one around the moment of maximum distortion, and one for the long-time limit. A strong change with time can be seen. The shift of the peak toward larger $I_1$ and smaller $I_2, I_3$ values already became clear from plots of the averaged values (Fig. 13).

The sharpest distributions are found around $t/\tau=O(10)$, where the peaks occur in the graphs of $I_i$. For long times, a considerable amount of tetrahedra has values for $I_1$ of the order of $0.5-0.6$, and values for $I_2$ of about $0.4-0.5$. This means that a reasonable amount of the tetrads (about 10%)
have a shape that is more two-dimensional than the average value of the eigenvalues would predict. $I_3$ is strongly peaked around zero, so few real three-dimensional structures exist.

Stable stratification suppresses vertical motion in the flow, resulting in less particle dispersion in vertical direction. When tetrahedra are placed in such a flow, it is expected that their shapes will become quasi-two-dimensional. The effect of stratification on the volume growth of a cluster of four particles can be seen in Fig. 15. With increasing stratification strength, the time at which tetrahedra start growing increases. Furthermore, it follows that with increasing stratification, the volume of a cloud decreases. Case N1000 seems to be an exception to this finding, but that is due to the chosen scaling with $\Delta_0$, which is well-suited for short times. When scaled with large-scale properties such as $T_L$ and $u_{\text{rms}}$, the order of the graphs for long times is as expected; the lowest graph then belongs to case N1000. Cases N1, N10, and N100 do not (yet) reach a final linear regime. For the strongest stratification studied in this work, the influence of initial cloud size on the growth of the cloud is shown in Fig. 16(a). As opposed to the isotropic case N0, for long times a difference can still be found here between different initial particle separations. The smallest three $\Delta_0$ reach similar cloud sizes, although their growth there is not equally strong. The smallest $\Delta_0$ does not (yet) show a final diffusive regime. The nondiffusiveness is presumably associated with the nondiffusive long-time horizontal single-particle dispersion.

The vertical direction, which is suppressed by stratification, plays a considerable role in the growth of a cloud. Even though in isotropic turbulence structures become mainly two-dimensional, the smaller volumes for stronger stratification point to even flatter structures in stably stratified turbulence. This is indeed the case, as can be concluded from the eigenvalues. For case N1000 in Fig. 16(b), the scaled eigenvalues $I_i$ are given as a function of time for different initial particle separations. It can be seen that $I_3 \approx 0$, $I_1 > (I_1)_{G4}$, and $I_2 < (I_2)_{G4}$. So the structures are more flattened than in the isotropic case. The time it takes to reach a final distribution is much longer than for case N0, and this distribution remains a bit unsteady. The influence of initial separation is similar to the case of isotropic turbulence. More generally, it can be concluded that with increasing $N$, $I_1$ increases, $I_2$ decreases, and $I_3$ converges to zero, resulting in flatter and more elongated structures. The long-time asymptotic values of the three eigenvalues are given in Table III for the different stratification levels. For all stratified cases, the long-time behavior of the eigenvalues is less stable than for case N0. Case N10, however, does not reach steady state within the duration of the simulation. The values given in the table for case N10 are therefore only rough estimates. The slight increase of $I_1$ and decrease of $I_2$ for case N1000 compared to case N100 can be caused by different initial particle separations ($\frac{5}{3} \eta$ and $\frac{5}{3} \eta$, respectively). The values for $I_1$ and $I_2$ can also be compared with the Gaussian values derived for multiparticle statistics in two-dimensional turbulent flow. Since the asymptotic value of $I_3$ is approximately zero for the strongly stratified cases, an option would be to compare the first two eigenvalues with the two eigenvalues that result from three-

![FIG. 15. Volume as a function of time for different strengths of the stratification: (N0: -- -- --, N1: - - - - - , N10: - - - - - , N100: - - - - - , N1000: - - - -). Volume is scaled by the initial particle separation $\Delta_0$. The choice of this scaling leads to the fact that the graph of case N1000 lies above that of case N100 for long times ($\Delta_{N1000}/\Delta_{N100}$). The inset shows the long-time behavior when scaled with $u_{\text{rms}}/T_L$, now cases N0 and N1 collapse and the order is as expected.](image1)

<table>
<thead>
<tr>
<th>Case</th>
<th>$I_1$</th>
<th>$I_2$</th>
<th>$I_3$</th>
</tr>
</thead>
<tbody>
<tr>
<td>N0</td>
<td>0.747±0.002</td>
<td>0.222±0.002</td>
<td>0.029±0.001</td>
</tr>
<tr>
<td>N1</td>
<td>0.808±0.006</td>
<td>0.179±0.004</td>
<td>0.014±0.004</td>
</tr>
<tr>
<td>N10</td>
<td>0.89±0.03</td>
<td>0.11±0.03</td>
<td>0.03</td>
</tr>
<tr>
<td>N100</td>
<td>0.856±0.004</td>
<td>0.143±0.004</td>
<td></td>
</tr>
<tr>
<td>N1000</td>
<td>0.841±0.003</td>
<td>0.158±0.004</td>
<td></td>
</tr>
</tbody>
</table>

![FIG. 16. For strongly stratified turbulence (case N1000), the volume (a) and eigenvalues $I_i$ (b) are shown as a function of time for different initial particle separations. Lines as in Fig. 13. These plots look a bit wobbly because of the limited amount of tetrahedron used to calculate statistics.](image2)
particle statistics in two-dimensional turbulence. This topic was studied by Castiglione and Pumir, who derived \( \langle I_2 \rangle_{G,3D} \approx 0.107 \) and thus \( \langle I_1 \rangle_{G,3D} \approx 0.893 \). The values found for four-particle statistics in stably stratified turbulence then lay in between those for four-particle statistics in isotropic 3D turbulence and three-particle statistics in 2D turbulence. Furthermore, we looked at three-particle statistics for case N100. Independent of the initial particle configuration, a purely horizontal or purely vertical triangle, the asymptotic value for \( I_2 \) is now 0.097. Comparing this value with the above-mentioned value for 2D turbulence \( \langle I_2 \rangle_{G,3D} = 0.107 \) and with the Gaussian result for three-particle statistics in 3D isotropic turbulence, \( \langle I_2 \rangle_{G,3D} = 1/6 \), our result resembles that for 2D turbulence.

For case N100, probability density functions of the eigenvalues \( I_1 \) are derived, and shown on the right-hand side of Fig. 14. All three eigenvalues show a peak around the time of maximum distortion that is much sharper than for case N0. Also for long times there is less spreading in the values of \( I_1 \). The distribution of eigenvalue \( I_1 \) is shifted toward 1 and that of eigenvalue \( I_2 \) towards zero. Eigenvalue \( I_3 \) only shows a very sharp peak close to zero. These values are a confirmation of the picture of flatter and more prolonged tetrahedra compared to isotropic turbulence. Moreover, it shows that the spectrum of structures occurring in stratified turbulence is less broad. For example, tetra with \( I_1 \approx I_2 \), denoting flat but not elongated shapes, are hardly found.

V. CONCLUDING REMARKS

Forced direct numerical simulations provide a means to study fluid particle dispersion in statistically stationary stably stratified turbulence. Stratification suppresses vertical fluid motion in turbulent flows. With increasing strength of the stratification, less and less small-scale motion and overturning is present in the flow. Suppression of vertical motion leads to the occurrence of plateaus in plots of both vertical single-particle and particle-pair dispersion. For long times, however, an increase in vertical dispersion is found approaching the linear diffusion limit as seen in isotropic turbulence. This regime is caused by molecular diffusion of the active scalar, the density \( \rho \), which slowly changes the equilibrium height of the particles. In horizontal direction, dispersion is enhanced for long times, especially for single-particle dispersion.

When looking at clusters of four particles, representing particle clouds, the method of deriving eigenvalues of the moment of inertialike tensor works well to describe shape evolution. Compared to isotropic turbulence, in stably stratified turbulence deformation of tetrahedra is enhanced. Structures tend to become more flattened and elongated and their volume decreases with increasing stratification. Moreover, the spreading in types of structures is much smaller than in isotropic turbulence.

The dispersion results presented in this work can be important for practical purposes. Both the suppression of vertical dispersion and the inhibition of cloud growth might enhance clustering and aggregation for interacting particles, because particles remain close together for longer times. On the other hand, the final diffusion limit found here makes it possible for particles to reach everywhere in the flow for long enough times.

To study dispersion in realistic geophysical environments, instead of passive fluid particles, particles with physical properties need to be studied. As a next step, we will include mass and inertial effects to study their influence on particle dispersion. Furthermore, the investigation of collisions and aggregate formation in stratified turbulence is left for future work.

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